

Carbonate Formation and Decomposition on Atomic Oxygen Precovered Au(111)

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The carbonate formation and decomposition $(CO_3 \leftrightarrow CO_2 + O_a)$ reaction on gold is important from the point of view of lowtemperature CO oxidation. Carbonate formation has been proposed as a possible reaction intermediate in CO oxidation in several investigations of supported and unsupported gold clusters.¹⁻⁴ Therefore, an understanding of this reaction on Au(111) may provide additional insight. Carbonate formation and decomposition went undetected in previous studies on Au(110)⁵ and Au(111).⁶ However, a surface carbonate was readily formed when oxygen precovered Ag(110) was exposed to CO₂ at 300 K.⁷⁻¹⁰ This surface carbonate decomposes to produce CO₂ at 485 K and the remaining oxygen atoms recombinatively desorbed at 590 K.⁷⁻¹⁰ Owing to its similarity with silver, we would anticipate equally facile carbonate formation and decomposition reactions on gold. Similar reactions have also been reported on other surfaces.¹¹⁻¹³

Here we present experimental evidence with supporting density functional theory (DFT) calculations of carbonate formation and decomposition from the adsorption of oxygen-labeled carbon dioxide ($C^{18}O_2$) on an atomic oxygen (^{16}O) precovered Au(111) surface. We studied the effects of CO₂ exposure, surface temperature, and oxygen coverage on carbonate formation and decomposition and also estimated reaction probabilities ($\sim 10^{-3} - 10^{-4}$) and activation energies as a function of conditions.

Our experiments were performed in a UHV chamber that has been described elsewhere,14-19 but details specific to this study are briefly summarized here. The Au(111) single crystal sample is mounted to a tantalum plate that can be resistively heated and is in thermal contact with a liquid nitrogen bath. Oxygen (¹⁶O) atoms were deposited using a radio frequency (RF) plasma-jet source. The ${}^{16}O_a/Au(111)$ surface was exposed to $C^{18}O_2$ by backfilling the chamber and carbonate ¹⁶OC¹⁸O¹⁸O was formed. The surface carbonate decomposes to form either C18O2 or 16O C18O leaving $^{18}O_a$ or $^{16}O_a$ adatoms on the surface. Upon heating, the oxygen atoms undergo recombinative desorption to produce ${}^{16}O_2$ (mass 32) and ¹⁶O¹⁸O (mass 34), as observed in TPD. Thus, carbonate formation and decomposition were detected via the increased presence of mass 34 18O16O in a temperature programmed desorption (TPD) spectrum after the ¹⁶O_a covered Au(111) surface was exposed to $C^{18}O_2$. We did not observe ${}^{18}O_2$ (mass 36) in TPD due to the very small surface concentration of ¹⁸O. This method was employed after other strategies proved unsuccessful because of the low reaction probability.

Figure 1a displays TPD spectra of ${}^{16}O^{18}O$ (m/e = 34) produced from exposure of the ${}^{16}O_a/Au(111)$ surface to $C^{18}O_2$. The amount of ${}^{16}O^{18}O$ produced increases with $C^{18}O_2$ exposure at 167 K (and all temperatures studied). Two control experiments were performed to ascertain the source of ${}^{16}O^{18}O$. First, no mass 34 was produced when the Au(111) surface was exposed to $C^{18}O_2$ without preadsorbed atomic oxygen. Second, only ~0.5% of the total amount



Figure 1. (a) TPD of ${}^{16}O{}^{18}O$ (*m/e* = 34) after a Au(111) surface covered with 1.3 ML ${}^{16}O$ at 77 K was exposed to varying amounts (0–30 L, where 1 L = 10^{-6} Torr s) of C ${}^{18}O_2$ at 167 K; (b) ${}^{16}O{}^{18}O$ production as a function of C ${}^{18}O_2$ exposure.



Figure 2. Integrated TPD area of ${}^{16}O{}^{18}O$ (mass 34) for varying initial oxygen precoverages (0.18–2.1 ML) on which 30 L of $C{}^{18}O_2$ was reacted at 167 K. The contribution to the signal due to the natural abundance of ${}^{18}O$ has been subtracted off.

of oxygen desorbs as mass 34 when the Au(111) surface is precovered with ¹⁶O but with no exposure to $C^{18}O_2$ (due to natural isotopic abundance of ¹⁸O). As expected, no surface-bound oxygen was lost during carbonate formation and decomposition, in agreement with previous studies.^{8–10} Figure 1b shows the amount of mass 34 produced (from the spectra in Figure 1a) as a function of $C^{18}O_2$ exposure.

To further examine the role of preadsorbed atomic oxygen on carbonate formation, we varied the oxygen precoverage (0.18–2.1 ML) while keeping both $C^{18}O_2$ exposure (30 L) and surface temperature (167 K) constant (Figure 2). Mass 34 production increases with increasing ${}^{16}O_a$ coverage, likely because more reactive oxygen is accessible to $C^{18}O_2$ on the surface. Similar results were obtained employing surface temperatures of 220 and 300 K.



Figure 3. Arrhenius plot of $C^{18}O_2$ reaction probability using a constant C¹⁸O₂ exposure of 30 L for 1.0 ML (upper plot) and 0.5 ML (lower plot) of atomic oxygen on Au(111).

We estimated the reaction probability of carbonate formation assuming a statistical distribution⁷ in the decomposition of the surface-bound carbonate 16OC18O2 and obtained values of ${\sim}10^{-3}{-}10^{-4}$ (uncertainties of $\pm50\%$). These small values are likely part of the reason why an earlier study on $Au(111)^6$ reported undetectable surface carbonate formation. An Arrhenius plot of the reaction probability for two oxygen coverages (0.5 and 1.0 ML) is shown in Figure 3. The inverse relationship between reaction probability and temperature, with negative apparent activation energy $E_a = -0.15 \pm 0.08$ eV, is suggestive of a competition between carbonate formation and C¹⁸O₂ desorption on the O/Au(111) surface.

Compared to Au(111), the carbonate formation reaction on Ag(110) is very facile.¹⁰ This difference is not currently understood but could be related to surface structure or other factors (likely not entirely due to the calculated energetics described below). Using DFT we have calculated the difference in energetics for CO₃ formation on Au(111), Au(110), Ag(111), and Ag(110). The metal surfaces were modeled with 4 (for 111) and 6 (for 110) layers, allowing the top two layers to relax. A vacuum gap of 10 Å separated the slabs. A plane wave basis set with a 274 eV cutoff was found to be sufficient for the PAW-based pseudopotentials,²⁰ with a $4 \times 4 \times 1$ Monkhost-Pack k-point sampling of the Brillouin zone. All calculations were based upon the PW91 GGA functional.21

Figure 4 shows calculated reaction paths for CO₃ formation on the metal surfaces. The initial point on each path corresponds to the binding of a single O atom at the most stable site on the surface, with the zero of energy taken with respect to gas-phase CO₂ and $1/2O_2$. CO₂ physisorbs to O_a and the two can react to form CO₃. The formation of CO_3 on Ag(110) was found to be spontaneous, and it occurs with only a very small barrier of 0.04 eV on Ag(111). For Au, however, there is a significant barrier to CO₃ formation, particularly on the (111) surface. Our calculations show that CO₃ is bound much more strongly to Ag than to Au, consistent with our experimental results on Au(111) in which CO₃ decomposition and CO₂ desorption appear to occur in an overlapping temperature range (90-120 K) while on Ag(110), the carbonate decomposes near 485 K and CO₂ desorbs below 160 K.

Our DFT calculations are consistent with the observed low reaction probability on Au(111) since CO₃ formation is activated as compared to CO₂ desorption. They do not explain the apparent negative activation energy; additional calculations including surface reconstructions may be necessary.



Figure 4. DFT calculations of carbonate formation on Au and Ag (111, upper plot) and (110, lower plot) surfaces. Energy barriers, ΔE_c , are labeled for each reaction pathway.

In summary, we have shown evidence for carbonate formation and reaction on atomic oxygen precovered Au(111). Oxygen mixing was observed when ¹⁶O_a precovered Au(111) was exposed to isotopically labeled CO_2 ($C^{18}O_2$) at surface temperatures ranging from 77-400 K and initial oxygen coverages ranging from 0.18-2.1 ML. Subsequent desorption of isotopically mixed oxygen (¹⁶O¹⁸O, mass 34) is observed as a byproduct of carbonate formation and decomposition on the surface. Carbonate formation occurs with a very small reaction probability (ca. $10^{-3}-10^{-4}$) and is most favorable at low surface temperatures.

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